

# Towards a simple NO<sub>2</sub> profile retrieval

MAX-DOAS  
KNMI

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Ankie Pitters  
Piet Stammes*

## Overview

MAX-DOAS TVCD retrieval based on AOT estimation from relative intensity (RI)

Choosing a reference for the RI (CINDI)

Towards a simple  $NO_2$  profile retrieval

Conclusions

DAK radiative transfer model comparison

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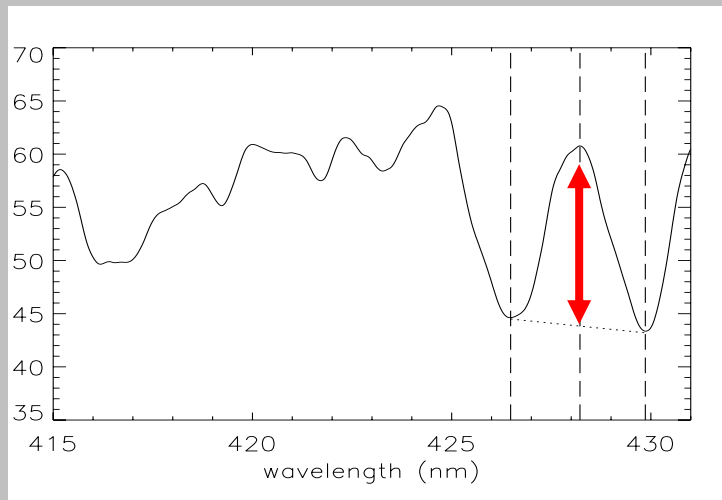
DAK radiative transfer model comparison

# MAX-DOAS differential air mass factors

→DAK radiative transfer model

→Atmosphere with NO<sub>2</sub> and aerosol in same layer (block profile)

→Direct calculation of DAMF from three wavelengths (similar to DOAS)

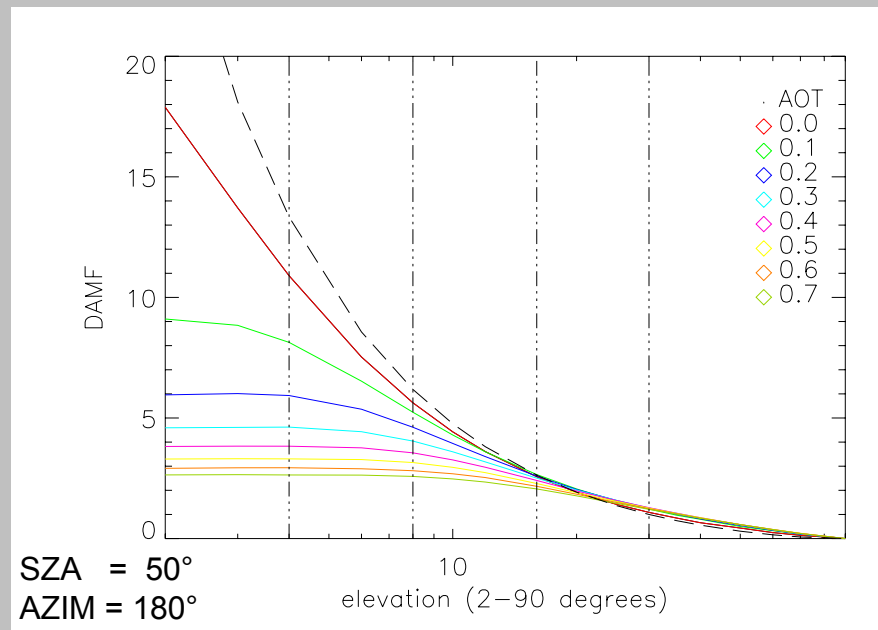


$$DAMF_{\alpha} = \frac{DSCD_{\alpha}}{TVCD}$$

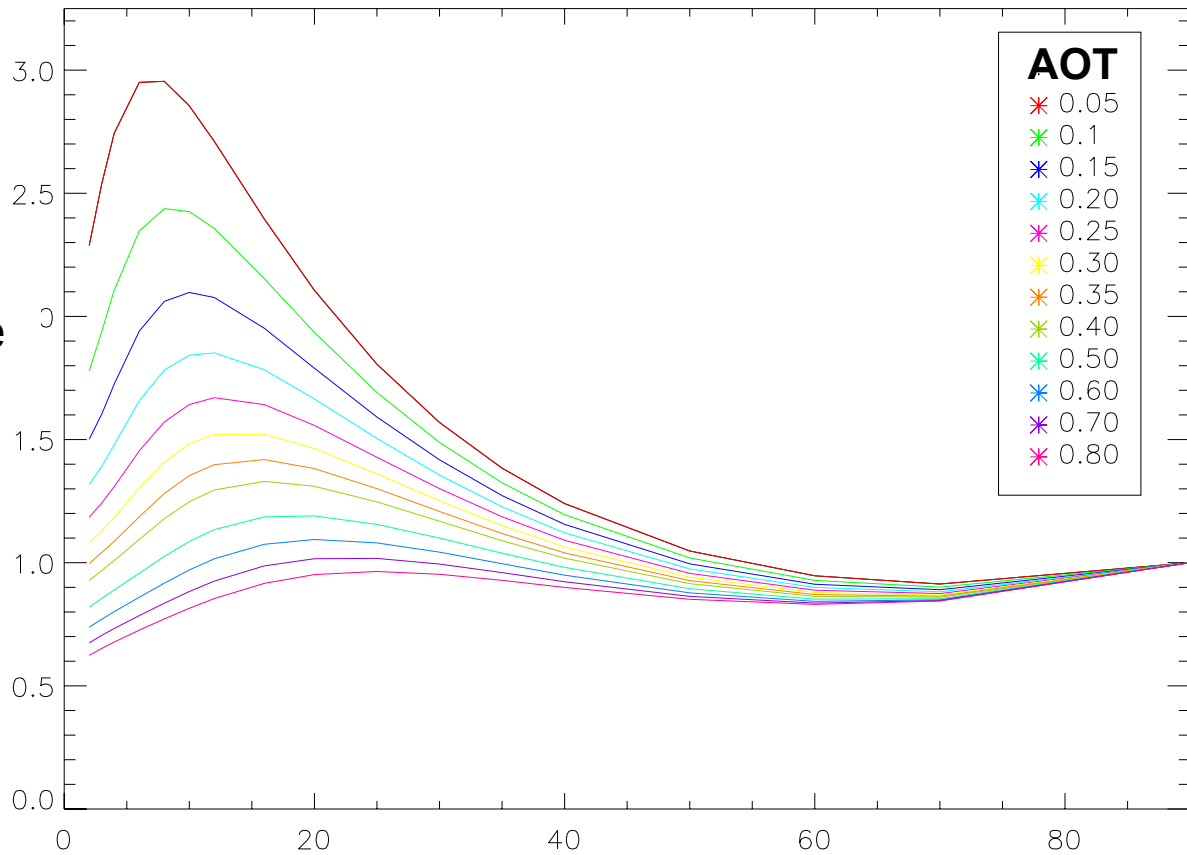
## Look-up tables: RI and DAMF

→ Depends on viewing geometry (sun, instrument), AOT, boundary layer height

→ Fixed parameters: profile shapes, asymmetry parameter, surface albedo etc.

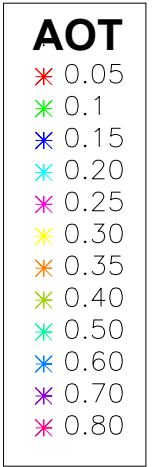


DAK simulation Azim: 140 SZA: 60 PBL: 1.0

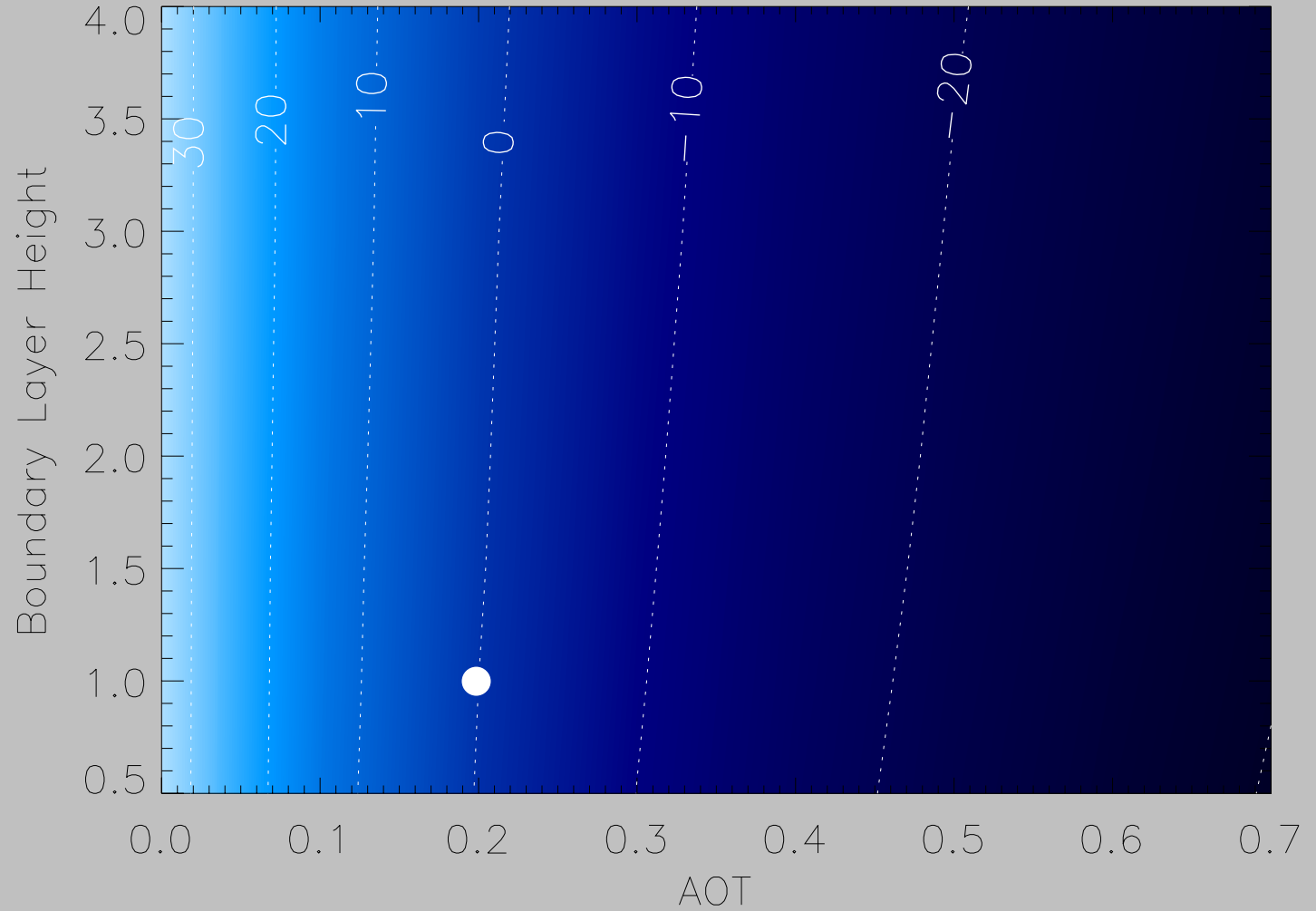


Intensity relative to zenith ( *RI* )

viewing angle (elevation)

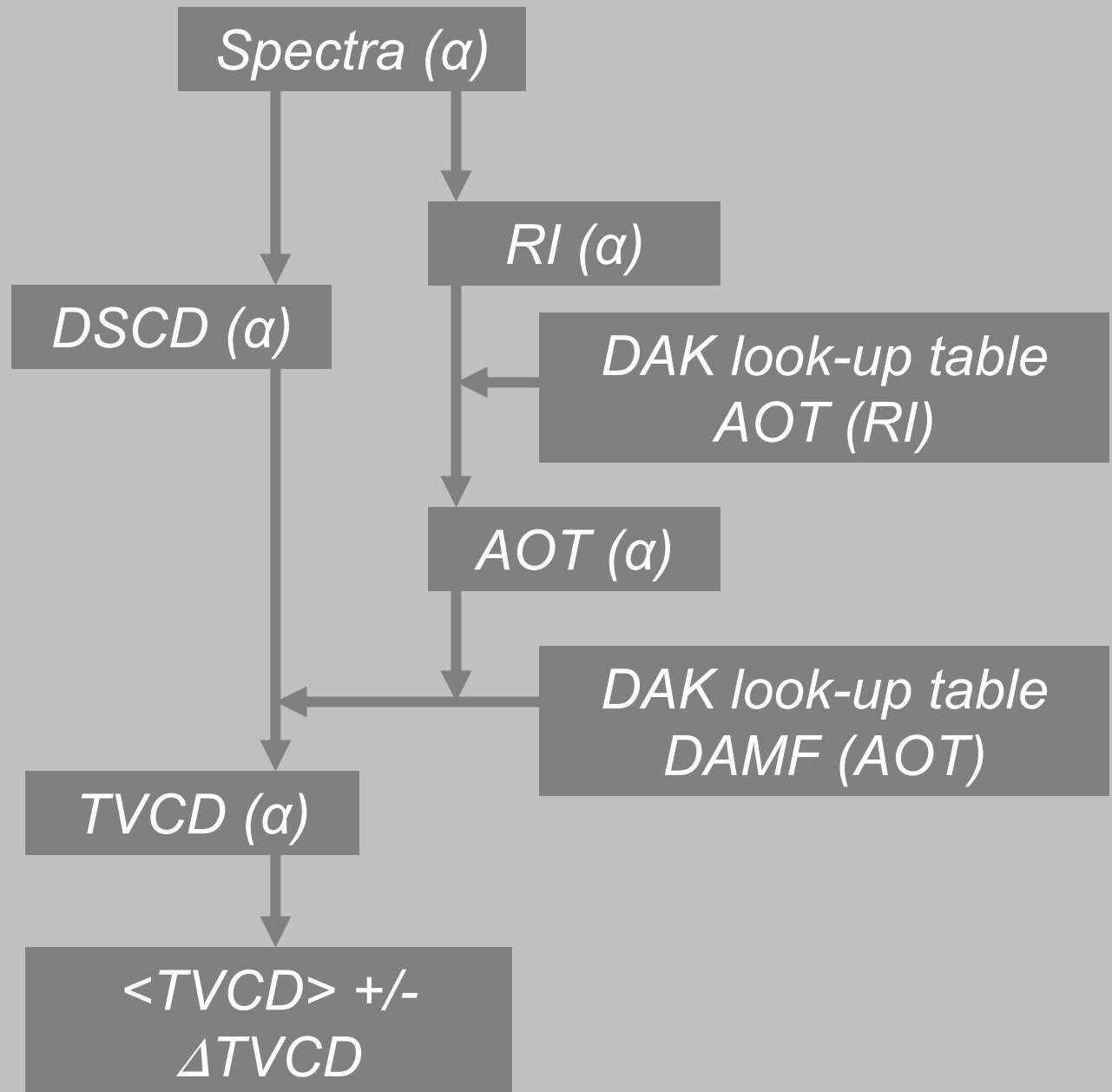


All values are % difference relative to: ●



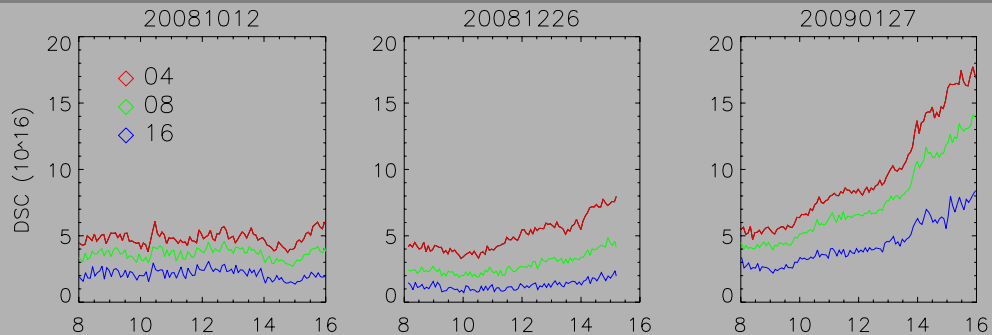
Relative Intensity @ 30°

# TVCD retrieval algorithm

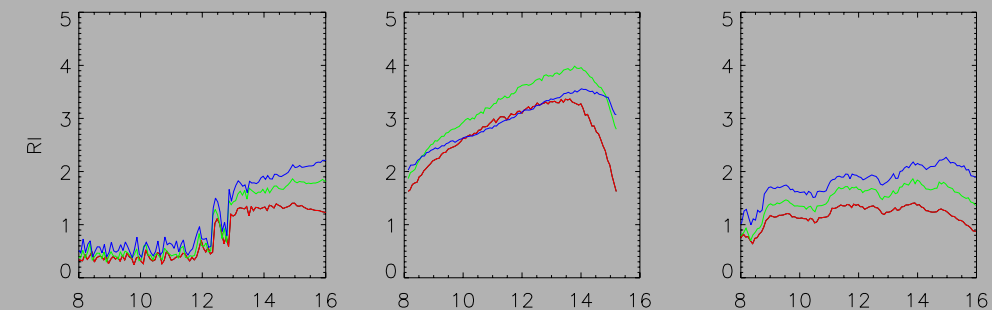




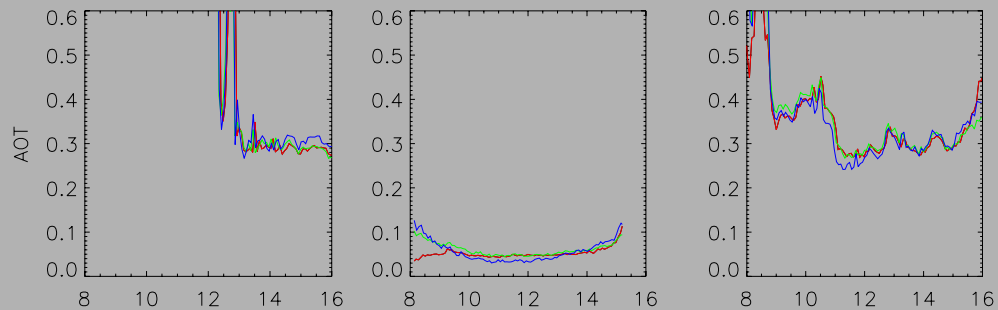
Differential  
Slant  
Column



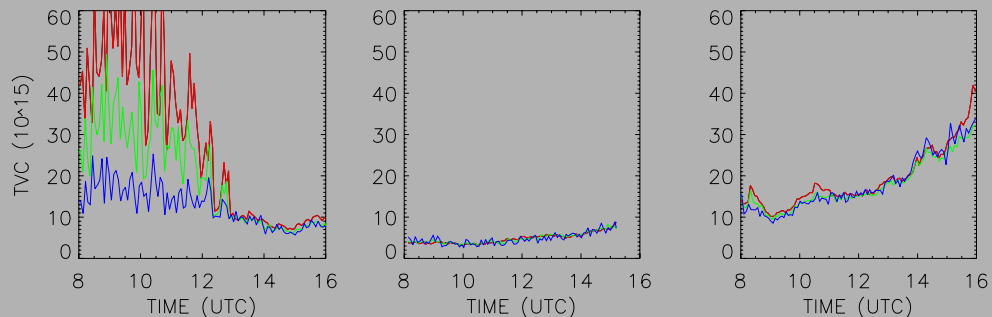
Relative  
Intensity

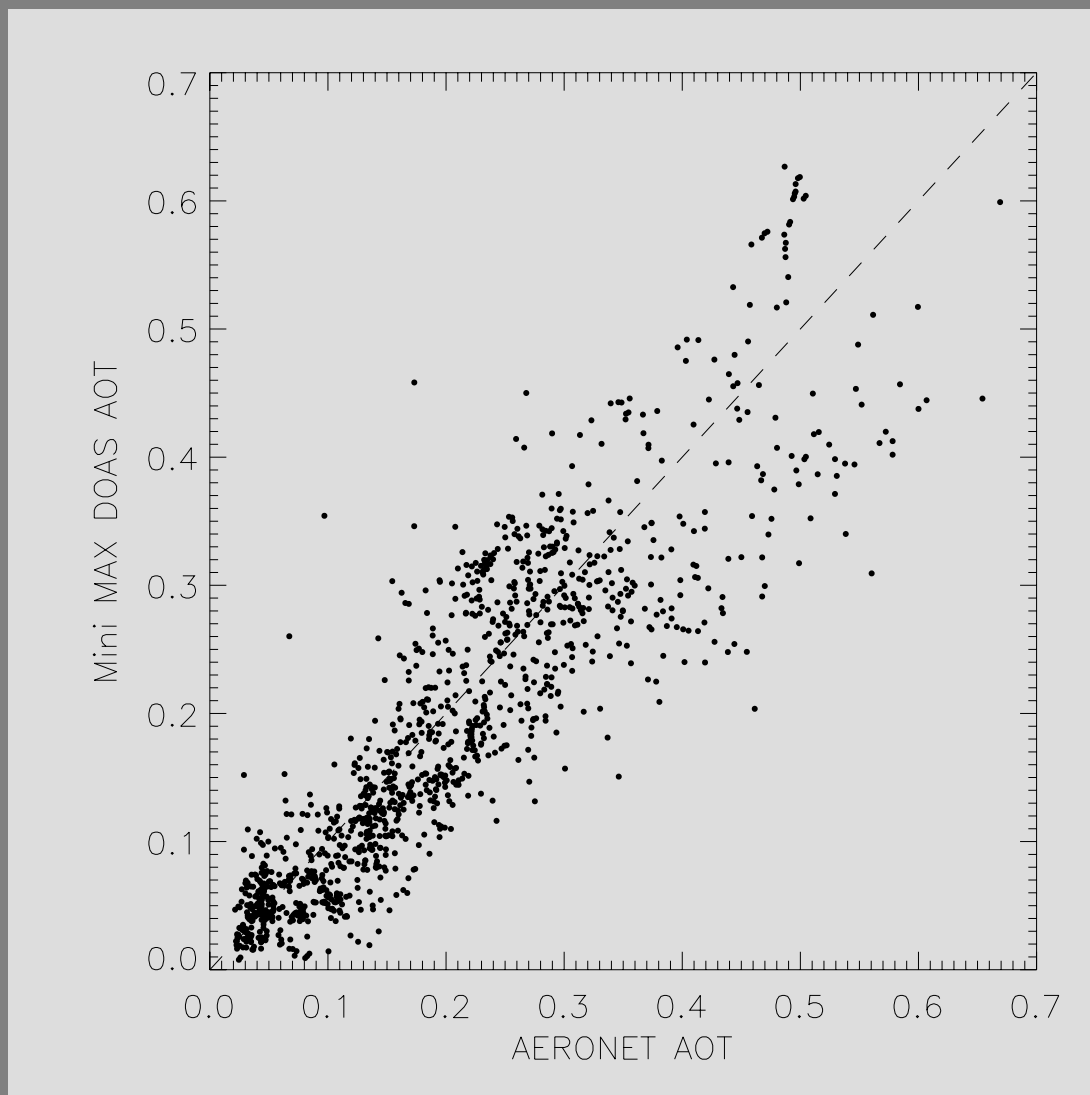


AOT



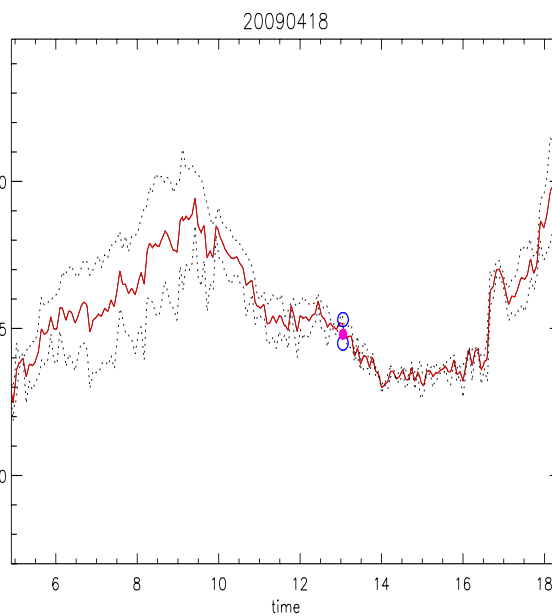
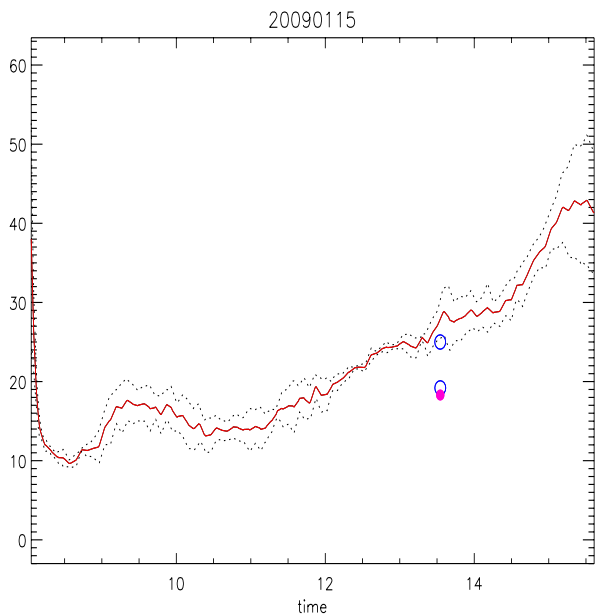
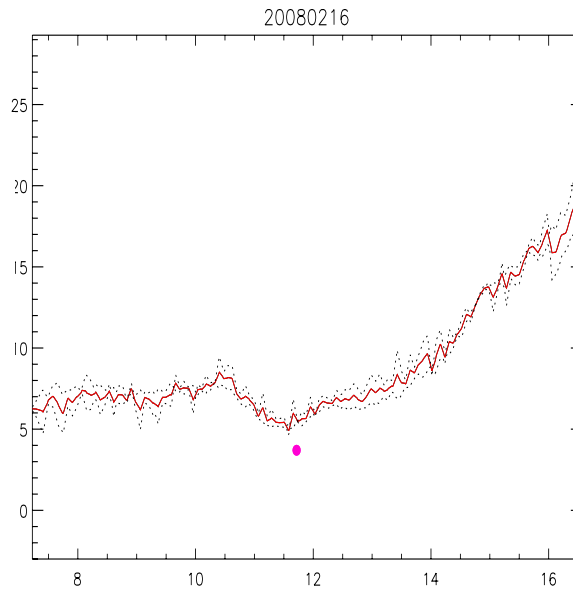
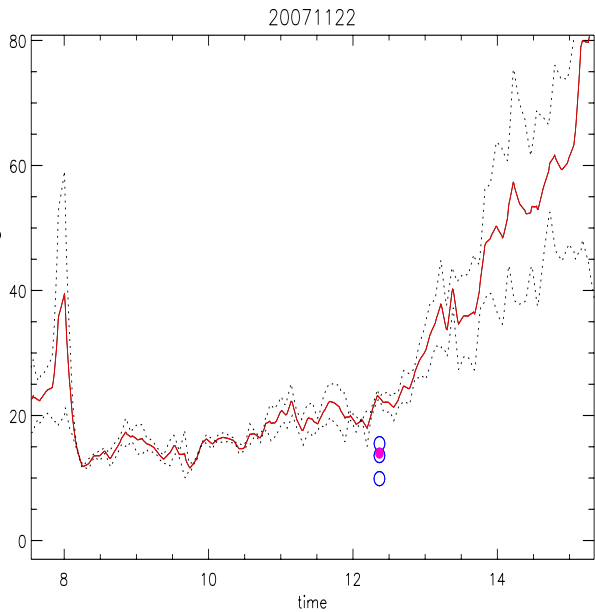
TVCD







# Tropospheric Columns: Example Days

MAX-DOAS TVCD  $\text{NO}_2 \cdot 10^{15} \text{ molec/cm}^2$



 TVCD averaged over 4°, 8°, 16° elevation

 Spread in TVCD derived from these elevations

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The RI-based retrieval was not so successful during the CINDI-campaign

Some possible reasons:

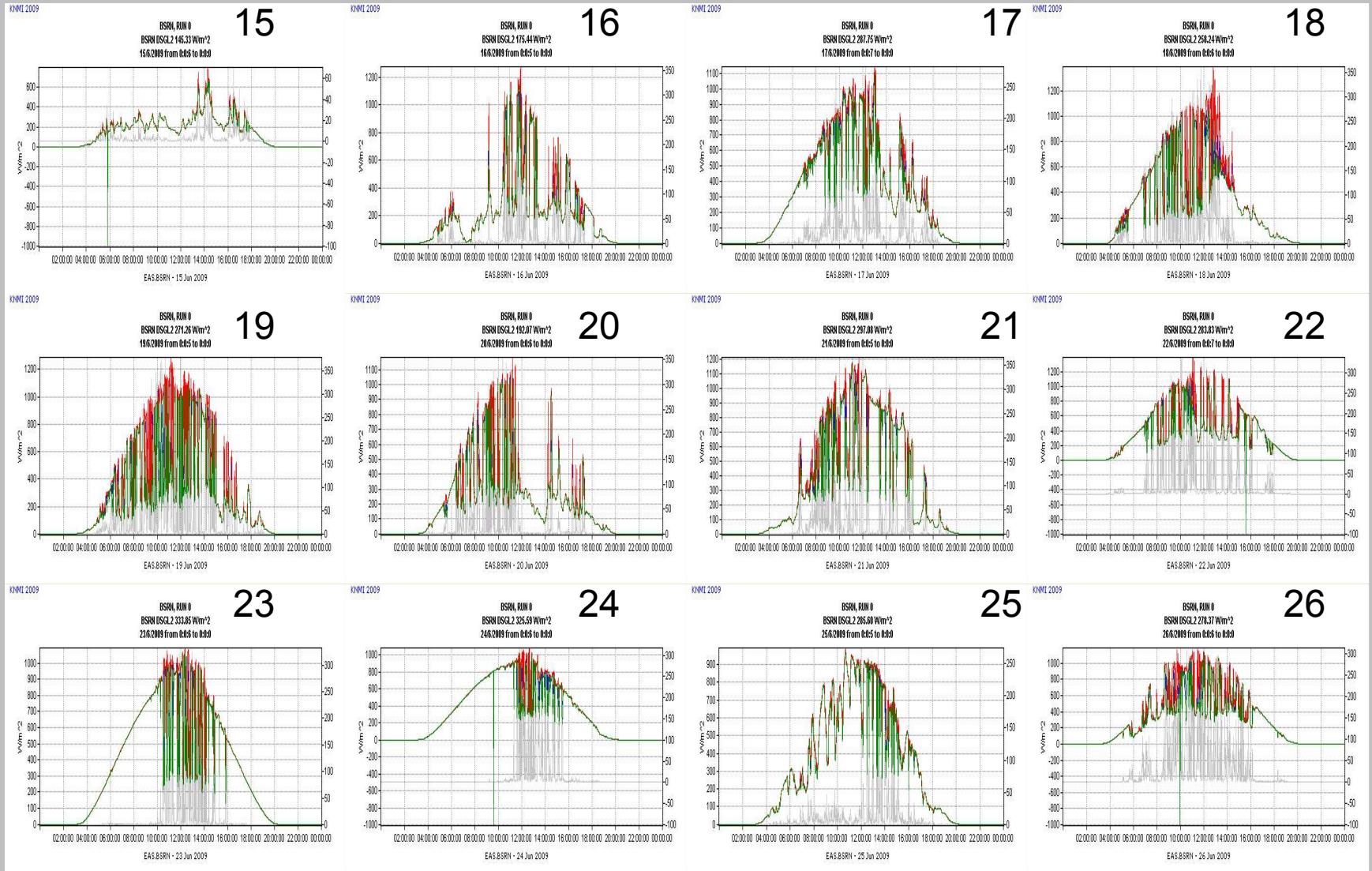
- many periods with clouds (homogeneous or scattered)

- the position of the sun

- we changed to longer integration times (2 minutes instead of 30 seconds per elevation which → no longer semi-simultaneous off axis and zenith contributions to RI)

# Direct and Diffuse radiation in Cabauw for CINDI semi blind period

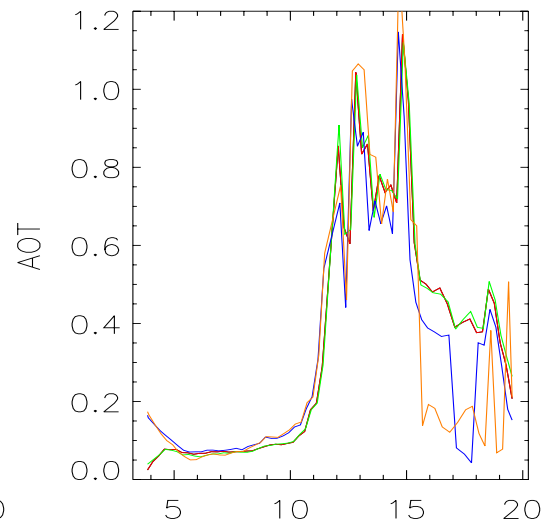
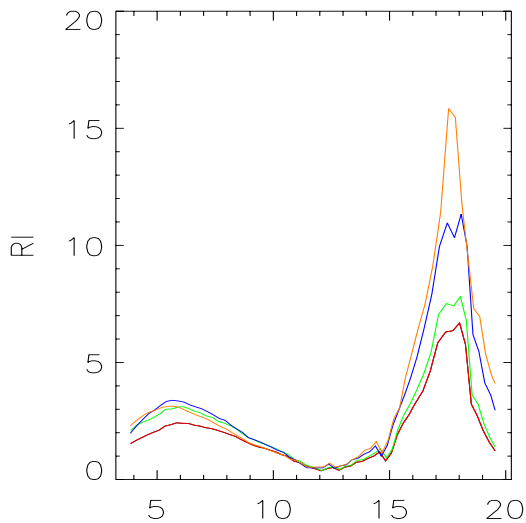
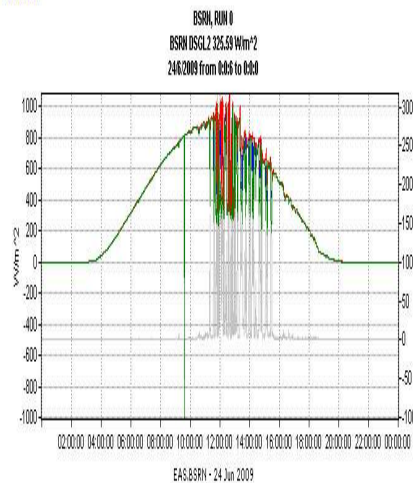
2009 / 06 / 15



2009 / 06 / 26

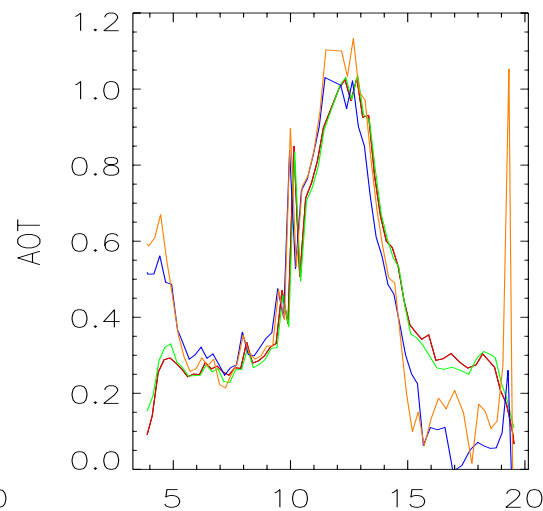
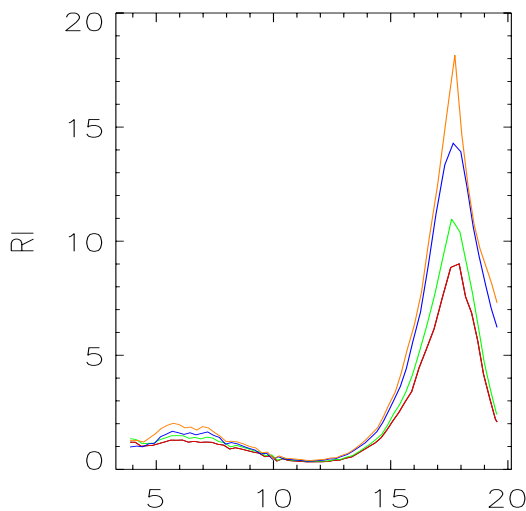
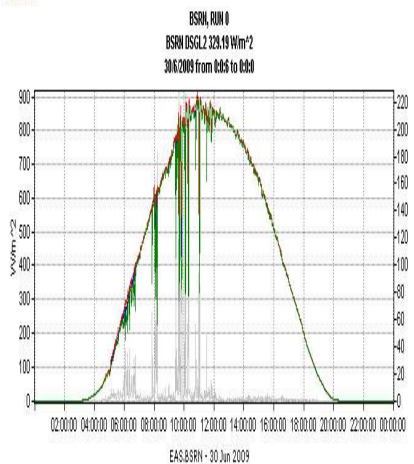
2009 / 06 / 24

KINZ 2009

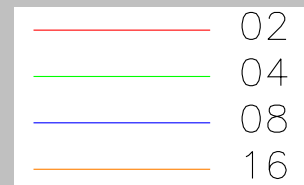


2009 / 06 / 30

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Elevation



CINDI: sun high in the sky around noon (SZA  $\sim 30^\circ$ );  
low sun in evening and close to the field of view

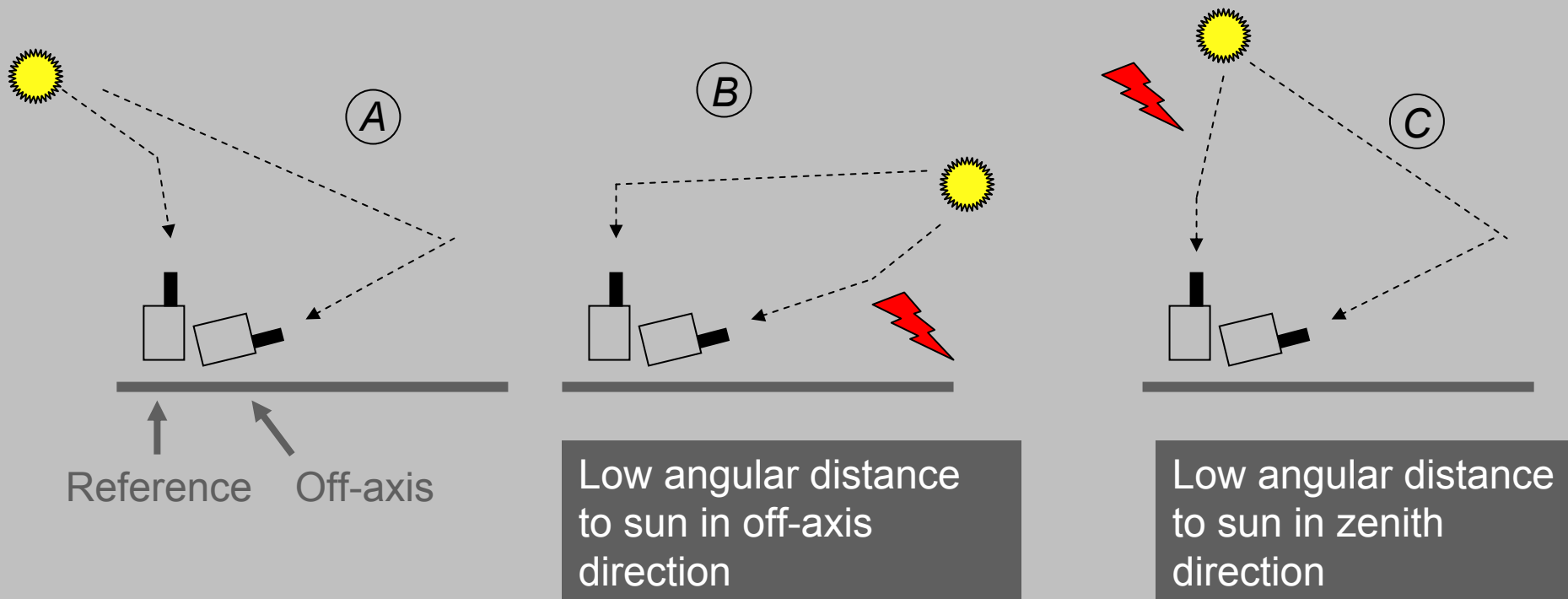
RT simulations of RI in viewing directions close to the sun depend more critically on the right choice of parameters in the model atmosphere

One of the main reasons is the shape of the aerosol phase function that is strongly peaked in a forward direction



RT simulations of RI in viewing directions close to the sun depend more critically on the right choice of parameters in the model atmosphere

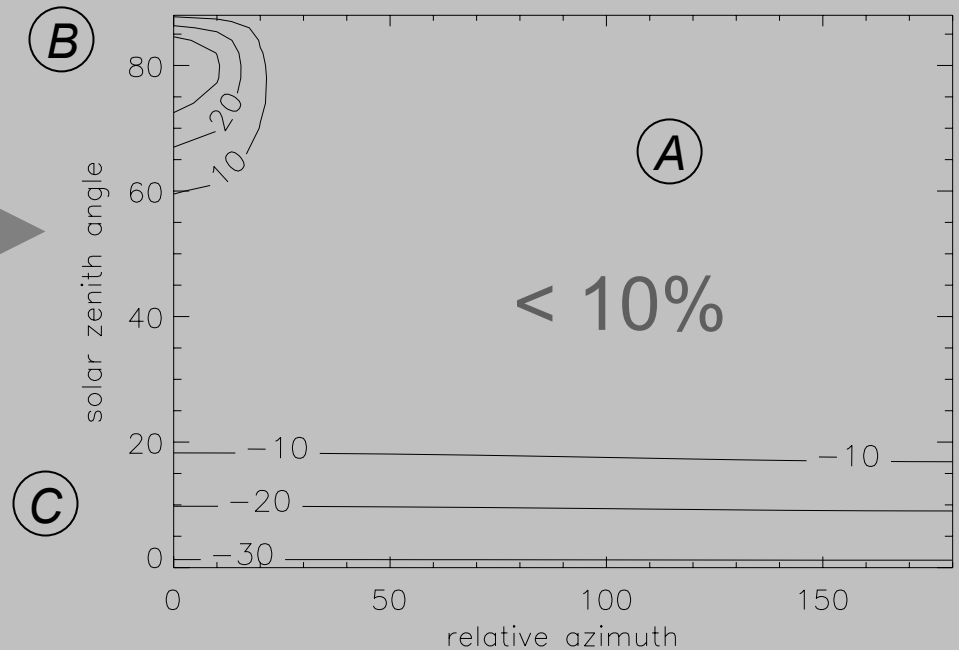
One of the main reasons is the shape of the aerosol phase function that is strongly peaked in a forward direction



# What does the DAK model say on this?

Change in RI (%) when  
**asymmetry parameter** is  
changed from 0.70 to 0.75 for  
elevation  $8^\circ$

$$RI = \frac{I_{8^\circ}}{I_{90^\circ}}$$

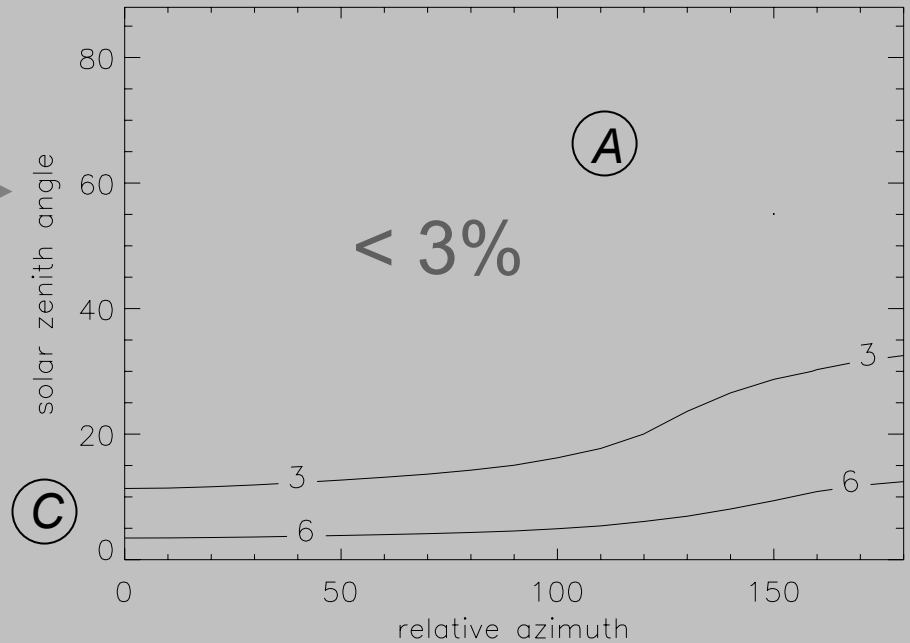


Inaccurate estimation of asymmetry parameter mainly affects  
measurements at low angular distance from the sun

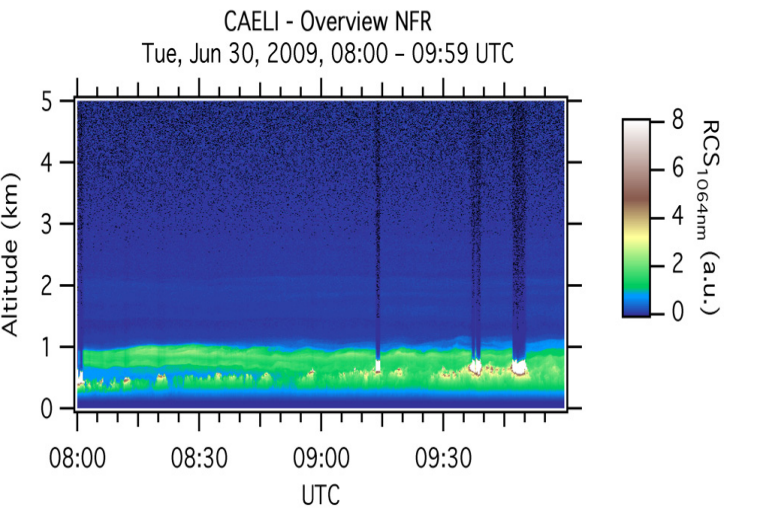
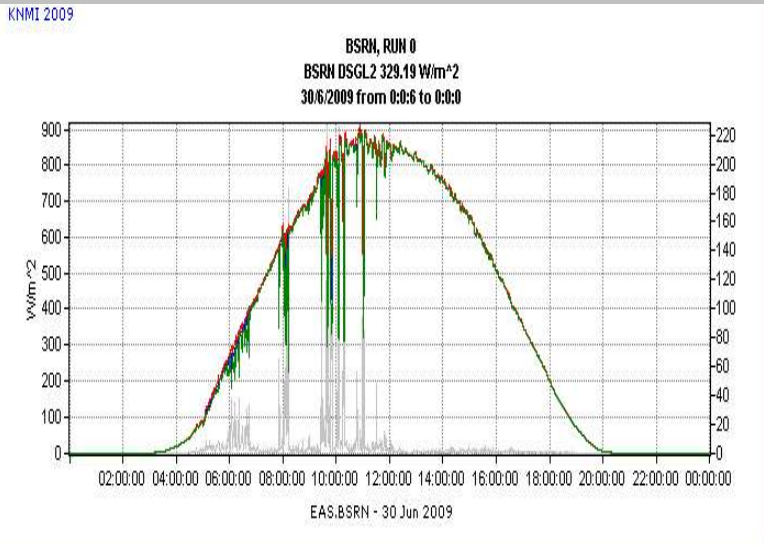
# A possible solution: change reference of RI

Relative Intensity now relative to 30°

$$RI = \frac{I_{8^\circ}}{I_{30^\circ}}$$



# Effect on the measurements



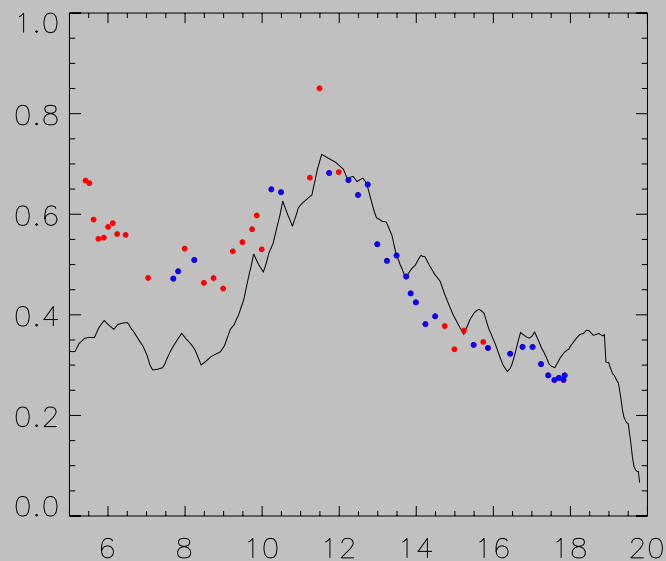
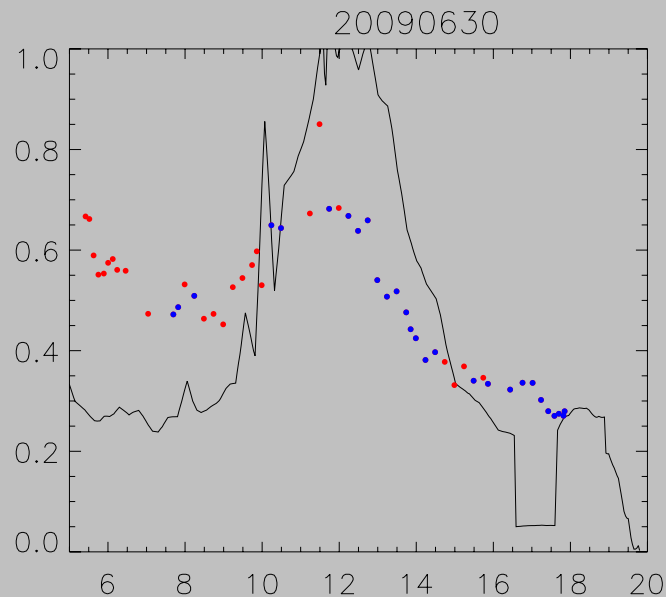
*Black line: AOT derived from RI*

$$RI = \frac{I_{4^\circ}}{I_{85^\circ}}$$

AOT

$$RI = \frac{I_{4^\circ}}{I_{30^\circ}}$$

AOT



● AERONET level 1.0

● AERONET level 1.5

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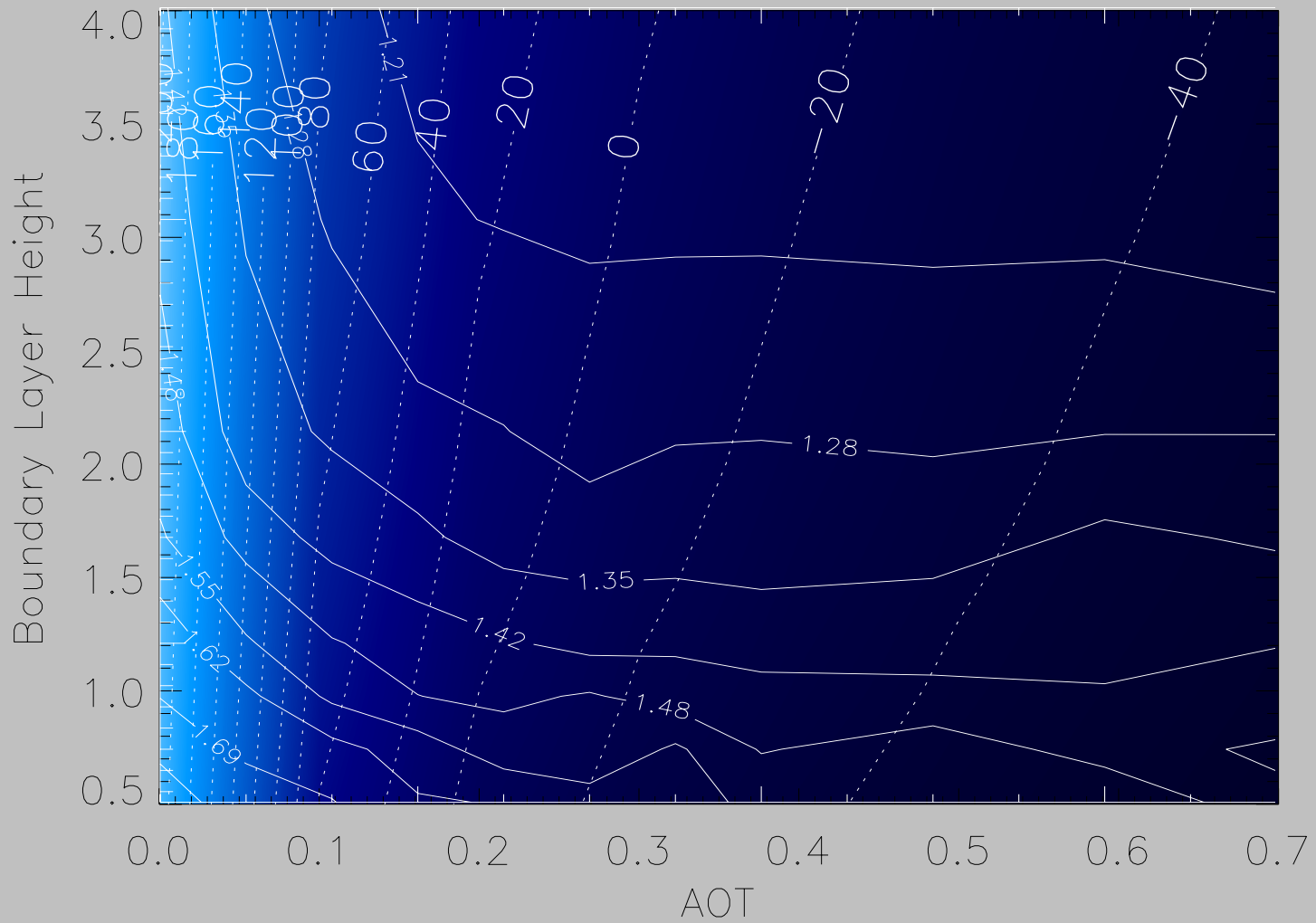
### General Approach:

Assume majority of NO<sub>2</sub> and aerosol in boundary layer (BL)

Retrieve only AOT and BL-height

Combine RI and O<sub>4</sub> absorption observations

Find orthogonal quantities in “AOT-BL space”



Example of orthogonal quantities in “AOT-BL space”

$$RI = \frac{I_{4^\circ}}{I_{90^\circ}} = \frac{DSCD_{-O_4}(4^\circ)}{DSCD_{-O_4}(16^\circ)}$$

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## Conclusions

In order to retrieve  $\text{NO}_2$  TVCD for MAX-DOAS an estimation of the AOT is needed

Relative Intensity (RI) provides a simple way to make a first order estimate of AOT in the BL

RTM simulations of RI require accurate knowledge of model parameters if observations are at close angular distance to the sun

- >> This limits the approach depending on fixed settings
- >> Change of reference may be a solution

A simple scheme is proposed to simultaneously derive AOT and BL-height from a combination of RI and  $\text{O}_4$  measurements

The simplified profile shape for  $\text{NO}_2$  is then a block function with the height of the BL

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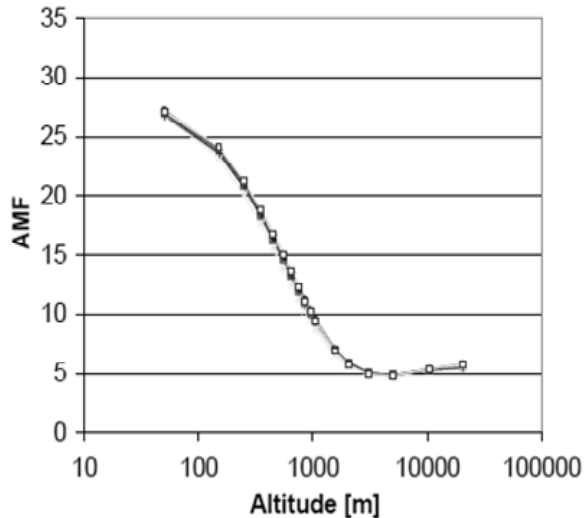
How does DAK RTM compare to other radiative transfer models?



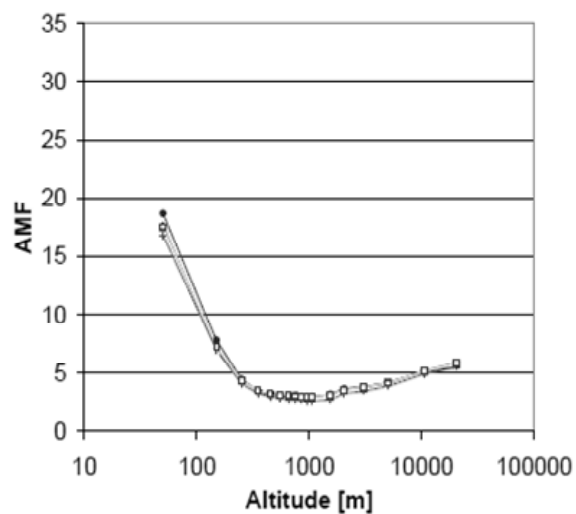
Wagner *et al.*, JGR 2007

Comparison of box-air-mass-factors and radiances for Multi-Axis Differential Optical Absorption Spectroscopy (MAX-DOAS) geometries calculated from different UV/visible radiative transfer models

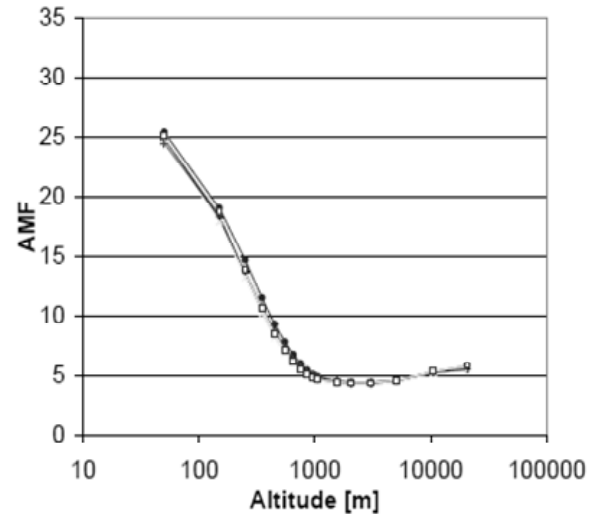
AMF, pure Rayleigh, 360 nm, 180° azim., 2° elev., 80° SZA



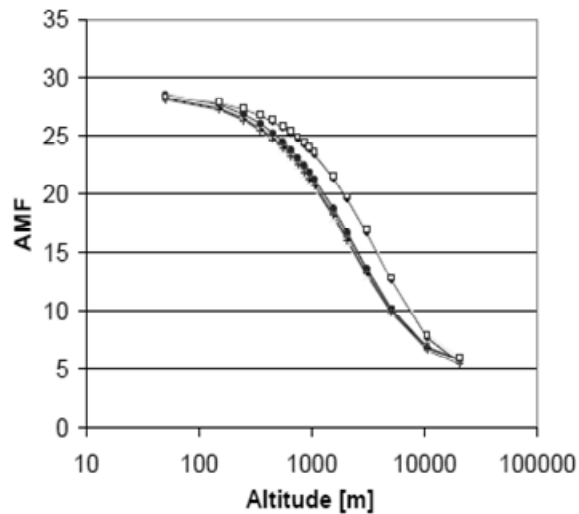
AMF, A2, 360 nm, 180° azim., 2° elev., 80° SZA



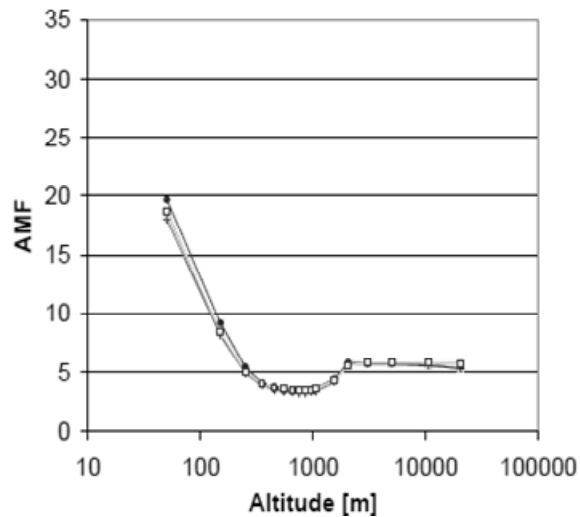
AMF, A3, 360 nm, 180° azim., 2° elev., 80° SZA



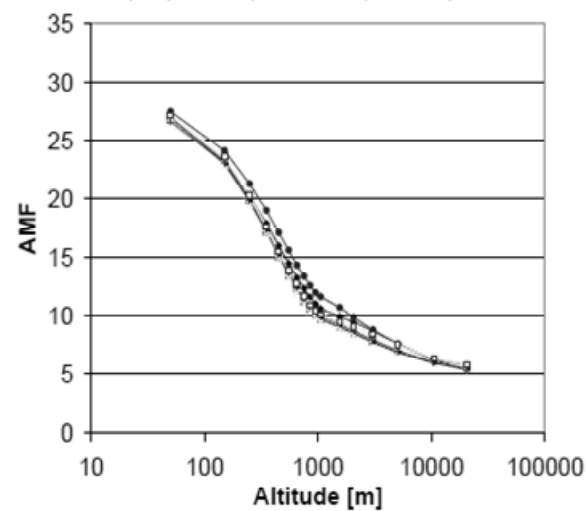
AMF, pure Rayleigh, 577 nm, 180° azim., 2° elev., 80° SZA



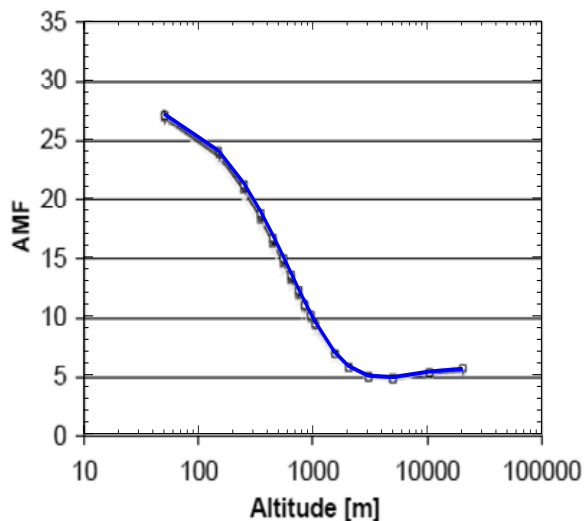
AMF, A2, 577 nm, 180° azim., 2° elev., 80° SZA



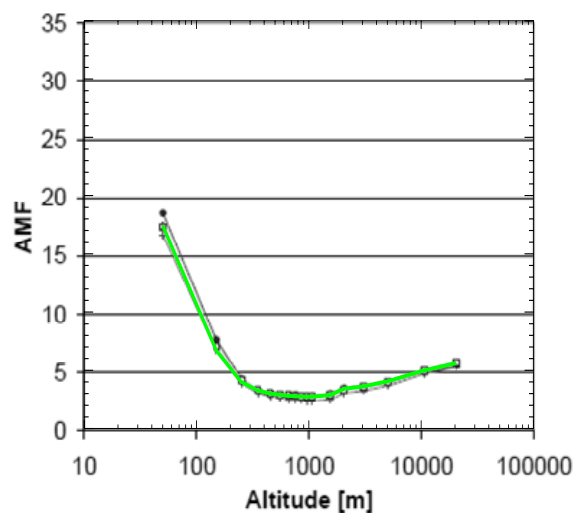
AMF, A3, 577 nm, 180° azim., 2° elev., 80° SZA



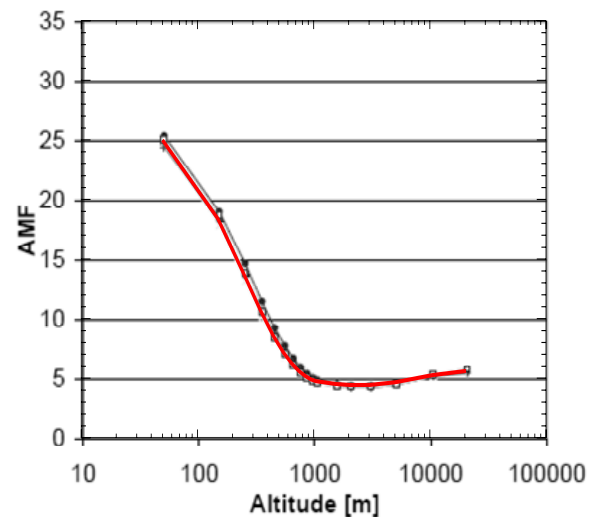
AMF, pure Rayleigh, 360 nm, 180° azim., 2° elev., 80° SZA



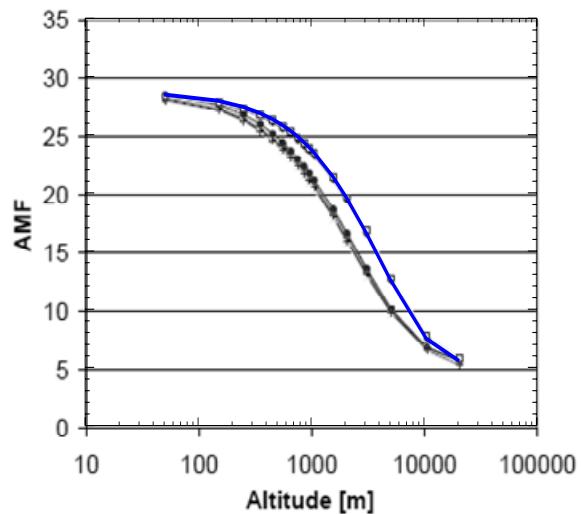
AMF, A2, 360 nm, 180° azim., 2° elev., 80° SZA



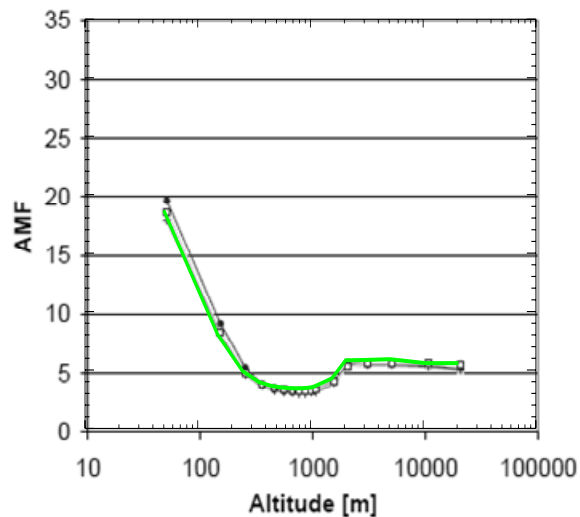
AMF, A3, 360 nm, 180° azim., 2° elev., 80° SZA



AMF, pure Rayleigh, 577 nm, 180° azim., 2° elev., 80° SZA



AMF, A2, 577 nm, 180° azim., 2° elev., 80° SZA



AMF, A3, 577 nm, 180° azim., 2° elev., 80° SZA

