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CINDI NO<sub>2</sub> workshop, Bremen, 18-19 Nov. 2009

## **BIRA's aerosol and NO<sub>2</sub> retrievals**





## Federal science pol



# The retrieval algorithm

# Forward model

Calculate radiances and  $O_4$  and  $NO_2$  DSCD @ different wavelengths and viewing geometries for a given atmosphere.

Linearized radiative transfer code (LIDORT v3.3) (R. Spurr, 2007)

INPUTS: - P, T

- Surface albedo
- Trace gases (σ,ρ)
- Aerosol (extinction profile single scattering albedo phase function)

Advantage of LIDORT : analytical calculation of weighting functions



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## Aerosol inversion

• Optimal Estimation method (Rodgers, 2000)

$$k_{i+1} = k_i + (S_a^{-1} + K_i^T S_{\varepsilon}^{-1} K_i)^{-1} [K_i^T S_{\varepsilon}^{-1} (y - F(k_i)) - S_a^{-1} (k_i - k_a)]$$

k = aerosol extinction vertical profile

 $k_a$  = apriori aerosol extinction vertical profile  $S_a$  = uncertainty covariance matrix of the apriori profile

F = Forward model (LIDORT) y = measurement ( $O_4$  DSCD and/or DI)  $S_{\epsilon}$  = uncertainty covariance matrix of the measurement

K = weighting functions =  $\partial y / \partial k$ 





# $NO_2$ inversion

Optimal Estimation method (Rodgers, 2000)

$$k = k_{a} + (S_{a}^{-1} + K^{T} S_{\varepsilon}^{-1} K)^{-1} K^{T} S_{\varepsilon}^{-1} (y - K k_{a})$$

 $\mathbf{k} = \mathbf{NO}_2$  vertical profile

 $k_a$ = apriori NO<sub>2</sub> vertical profile S<sub>a</sub> = uncertainty covariance matrix of the apriori profile

F = Forward model (LIDORT)y = measurement (NO<sub>2</sub> DSCD) $S<sub>\varepsilon</sub> = uncertainty covariance matrix of the measurement$ 

K = weighting functions = 
$$\partial y / \partial \mathbf{k}$$

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# Flowchart



### CINDI WORKSHOP, KNMI, 6-8 July 2009



# BIRA settings: O<sub>4</sub> xs

## **Based on the Beijing dataset Case**: 30° elevation , pointing north, clear-sky , AOD<0.15



## Measured and simulated $O_4$ DSCD should be equal



## But sim. $O_4$ DSCDs = meas. $O_4$ DSCDs \* 0.8±0.1



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## **BIRA** settings

FORWARD MODEL

P,T: Radio sondes
 surface albedo: lambertian = 0.07
 O<sub>3</sub> US standard profiles

□ Single scattering albedo and phase function calculated using a Mie routine and inputs from the AERONET.

## **OPTIMAL ESTIMATION**

O<sub>4</sub> DSCD \* 0.75
O<sub>4</sub> apriori: exponential profile; 1 km scaling height; AOD=0.05
Sɛ: diagonal, (DSCD error)<sup>2</sup>
S<sub>a</sub>: see next slide
Only retrieve 4 km.
non-linear equation for optimal estimation

## $\square$ NO<sub>2</sub> DSCD

□ apriori: NO<sub>2</sub> apriori: US standard + 0.25 ppb in the lowest layer, 0.05 ppb at 4 km and a linear decrease in between.

- $\Box$  Se: diagonal, (DSCD error)<sup>2</sup>
- $\Box$   $S_a$ : see next slide
- $\Box$  Only retrieve 4 km.
- □ Linear equation for OE





# BIRA settings: $S_a$

 $\Box S_a$  changes each iteration

**L**owest layer:  $S_a(1,1)=(factor*maximum(\mathbf{x}_i))^2$ 

 $\Box$ At 4 km:  $S_a(n,n)=0.2*S_a(1,1)$ 

In between a linear decrease with altitude

□Off-diagonal elements were set using Gaussian correlation functions with a correlation length of 0.05km.

## $NO_2$ :

Aerosol:

 $\Box S_a = (factor*apriori)^2$ 

□Correlation length = 0.2 km



 $\square For aerosol and NO_2$  retrieval the factor making up the  $S_a$  is chosen to have a mean DFS of ~2.

Factor aerosol = 0.1 for BIRA and Bremen data
 Factor NO<sub>2</sub> = 0.8 for BIRA data
 = 0.3 for Bremen data











# Results: aerosol

## **Results:** Aerosol





## .







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## aerosol surface extinction





Paul Zieger; PSI Switserland

# aerosol: Bremen data







## aerosol: compare



MAXDOAS Bremen data 100m [1/km] 477nm

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# NO<sub>2</sub> surface concentration









## NO<sub>2</sub> VC aerosol impact







# NO<sub>2</sub> surface concentration

## aerosol impact







# NO<sub>2</sub> surface concentration











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## Summary

Retrieval based on LIDORT and OE.
 Correction factor for O<sub>4</sub> cross section xs=xs/0.75.
 S<sub>a</sub> for aerosol retrieval changes each iteration.
 S<sub>a</sub> tuned so that the DFS~2.

 $\Box$ Aerosol have substantial influence on NO<sub>2</sub> profile shape.

□It all looks quite promising.

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## **MAXDOAS** measurement

## Multi-AXis Differential Optical Absorption Spectroscopy



 Increased sensitivity towards atmospheric absorbers present close to the surface
 eliminate strato. contr. □ Use high frequency differential absorption structures to identify absorbers and quantify their abundance

Atmosphere I absorption Aerosol / Molecules I Remote sensing instrument

DOAS





□ Provides information on the **vertical distribution** of gases and aerosol in the troposphere.



ISTP 2009, Delft, 20-Oct-09

## **UV-VIS Channels**



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